

1 **Dark aerobic sulfide oxidation by anoxygenic phototrophs in the anoxic waters**  
2 **of Lake Cadagno**

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28 **ORIGINALITY-SIGNIFICANCE STATEMENT**

29 This study reveals that sulfide oxidation within an anoxic layer of purple sulfur bacteria in the stratified  
30 water column of Lake Cadagno is largely coupled to oxygen consumption. Our findings imply that aerobic  
31 metabolism may be more prevalent in anoxic zones than previously thought. We also present a  
32 metagenome-assembled genome of *Chromatium okenii* which is the first genome sequence for the genus  
33 *Chromatium* and reveals new interesting physiological features of this environmentally relevant organism  
34 including its capacity for aerobic respiration.

35 **SUMMARY**

36 Anoxygenic phototrophic sulfide oxidation by purple and green sulfur bacteria plays a key role in  
37 sulfide removal from anoxic shallow sediments and stratified waters. Although some purple sulfur  
38 bacteria can also oxidize sulfide with nitrate and oxygen, little is known about the prevalence of  
39 this chemolithotrophic lifestyle in the environment. In this study, we investigated the role of  
40 *Chromatium okenii* in chemolithotrophic sulfide removal in the chemocline of Lake Cadagno. This  
41 purple sulfur bacterium appears to remain active during the night, as evidenced by its continued  
42 motility and O<sub>2</sub>-driven carbon fixation. Our temporally resolved, high-resolution chemical profiles  
43 revealed that sulfide oxidation is largely driven by aerobic respiration in the anoxic chemocline.  
44 We postulate that the abundant and highly active *Chr. okenii* are, at least in part, responsible for  
45 this aerobic sulfide oxidation and that they bridge the spatially separated gradients of oxygen and  
46 sulfide using a novel mechanism of transport driven by the strong convection within the  
47 chemocline. The genome of *Chr. okenii* reconstructed from the Lake Cadagno metagenome  
48 confirms its capacity for microaerophilic growth and provides further insights into its metabolic  
49 capabilities. Altogether, our observations suggest that aerobic respiration may not only play an

50 underappreciated role in anoxic environments, but also that organisms typically considered strict  
51 anaerobes may be involved.

## 52 **INTRODUCTION**

53 Anoxygenic phototrophic bacteria oxidizing sulfide and fixing CO<sub>2</sub> with sunlight play an important  
54 role in the carbon and sulfur cycles of sulfidic, shallow sediments and stratified water columns.  
55 Phototrophic sulfur bacteria, for example, are responsible for 20-85% of the total daily carbon  
56 fixation in anoxic lakes (summarized in Cohen *et al.*, 1977). This primary production is so  
57 important that it can control the bulk C-isotope fractionation in the water column, generating  
58 isotopic signatures that are transported and preserved in sediments (Posth *et al.*, 2017). Biomass  
59 from anoxygenic phototrophs feeds both grazing zooplankton in overlying oxic waters (Sorokin  
60 1966) and drives sulfate reduction in anoxic waters below (Pfennig 1975). The phototrophic sulfur  
61 bacteria also remove toxic sulfide from the water column enabling aerobic life at the surface while  
62 recycling sulfur compounds for sulfate reducers. While their role in sulfide detoxification has long  
63 been recognized in stratified lakes, there is mounting evidence that phototrophic sulfur bacteria  
64 also significantly impact sulfur cycling in marine environments such as the Black Sea (Jørgensen  
65 *et al.*, 1991) and the Chesapeake Bay (Findlay *et al.*, 2015).

66 Anoxygenic phototrophs generally inhabit illuminated, anoxic, reducing environments due to the  
67 toxicity of oxygen to these bacteria, and to the competition with abiotic reactions involving oxygen  
68 for their electron donors. Nonetheless, some anoxygenic phototrophs have evolved the capacity  
69 for chemotrophic growth under microoxic conditions. Whereas the green sulfur bacteria (GSB) of  
70 the *Chlorobiaceae* family are strict anaerobes, members of *Proteobacteria* collectively known as  
71 the purple sulfur bacteria (PSB), can be anaerobic to microaerobic (e.g. Kampf and Pfennig, 1980;

72 de Witt and Van Gernerden, 1990). Both the GSB and PSB are well adapted to fluctuating  
73 environmental conditions, synthesizing and accumulating storage compounds during periods of  
74 nutrient excess. The anoxygenic phototrophs are known to store zero-valent sulfur ( $S^0$ ),  
75 polyphosphate, glycogen, and in the case of the PSB alone, poly-3-hydroxyalkanoates (PHA) (Mas  
76 and Van Gernerden, 1995). The macromolecular structure and metabolism of these compounds  
77 have been intensely studied in laboratory pure cultures in order to understand conditions leading  
78 to their accumulation and breakdown. It has been suggested that glycogen may play a role in  
79 energy generation under dark conditions based on observations that cultured *Chromatium* sp.  
80 utilize glycogen to reduce stored sulfur, yielding sulfide and PHA (Van Gernerden, 1968).

81 Here we investigated the role of anoxygenic phototrophic bacteria in dark sulfur cycling processes  
82 in Lake Cadagno, a permanently stratified lake with high sulfate concentrations of up to 1-2 mM  
83 in the monolimnion. Microbial reduction of sulfate in the anoxic bottom waters and sediments  
84 produces large amounts of sulfide which support dense populations of GSB and PSB in the photic  
85 zone. These bacteria heavily influence the chemistry of the lake, forming a sulfide- and oxygen-  
86 free chemocline of 1-2 meters in thickness. The PSB *Chromatium okenii* is by far the most active  
87 of these bacteria, having been shown to play a disproportionately large role in inorganic carbon  
88 and ammonium assimilation despite their low abundances (<1% of total cell numbers) in the  
89 chemocline (Musat *et al.*, 2008; Posth *et al.*, 2017). In addition to their important contribution to  
90 light-driven sulfide oxidation, previous studies have shown that the anoxygenic phototrophic  
91 bacteria of Lake Cadagno remain active in the dark (Musat *et al.*, 2008; Halm *et al.*, 2009; Storelli  
92 *et al.*, 2013). However, their mechanism of energy generation in the absence of light is not yet  
93 clear. There is also evidence for dark sulfide consumption, but the electron acceptors utilized

94 remain unknown. We therefore combined high-resolution biogeochemical profiling with  
95 metagenomic analyses to gain an overview of possible light-independent metabolic processes  
96 impacting the sulfur biogeochemistry of Lake Cadagno. In addition to providing insights into the  
97 metabolism of anoxygenic phototrophic bacteria *in situ*, we present a model to explain the  
98 mechanism of dark sulfide oxidation in the chemocline of this meromictic lake.

## 99 **RESULTS & DISCUSSION**

### 100 ***Biogeochemistry of Lake Cadagno***

101 Lake Cadagno is characterized by an oxic mixolimnion and a sulfidic monolimnion spatially  
102 separated from each other by a chemocline (defined by bold contour lines in Fig. 1a) free of  
103 detectable oxygen (detection limit 50-100 nM) and containing very little sulfide. In August 2015,  
104 oxygen disappeared just above the chemocline close to 12 m depth. The daytime increase in  
105 oxygen concentrations between 11-12 m depth denotes net photosynthesis and the nighttime  
106 decrease denotes net respiration (Fig. 1a). The permanent absence of oxygen in the chemocline  
107 indicated that oxygen was consumed both in the day and the night.

108 Steep gradients of sulfide diffusing into the chemocline varied independently of light-dark periods  
109 and the total sulfide concentration in the chemocline did not exceed 5  $\mu\text{M}$  at any time point.  
110 Because the lake is meromictic, these stratified conditions were also present during other  
111 sampling years (see Fig. S2 for 2013 and 2014 profiles). In 2015, the 0.5-1 m wide chemocline was  
112 located around 11-12 m depth, with the exact location varying over the day most likely due to the  
113 action of internal waves (Egli *et al.*, 1998). In previous years, the chemocline was up to 2 m wide  
114 (Fig. S2) and remained completely sulfide-free in the dark. Conservative properties such as  
115 temperature and conductivity were constant throughout the chemocline in all years sampled (Fig.

116 S1&2) indicating mixing of this zone (Sommer *et al.*, 2017). Flat conductivity profiles revealed  
117 stronger mixing of the chemocline in 2013 and 2014 (Fig. S2) than in 2015 (Fig. S1) when the  
118 region of constant conductivity was reduced or absent.

119 *Chr. okenii* was the most significant microorganism in the chemocline in terms of biomass,  
120 accounting for ~60-80% of total microbial biovolume (Sommer *et al.* 2017), and carbon fixation  
121 (Musat *et al.*, 2008). The cell abundances of *Chr. okenii* in the Lake Cadagno chemocline were  
122 enumerated by flow cytometry during 2 daily cycles (Fig. 1b). Higher densities of *Chr. okenii* were  
123 found in 2014 ( $10^6 \cdot \text{ml}^{-1}$ ) than in 2015 ( $10^5 \cdot \text{ml}^{-1}$ ). *Chr. okenii* is highly motile, swimming at speeds  
124 of  $\sim 27 \mu\text{m} \cdot \text{s}^{-1}$  and has been hypothesized to drive the convection and mixing of the chemocline  
125 (Wüest, 1994; Sommer *et al.*, 2017). *Chromatium* are known to migrate between gradients of  
126 sulfide, light, and oxygen by photo- and chemotaxis (Pfennig *et al.*, 1968). We observed that *Chr.*  
127 *okenii* were positioned between oxygen and sulfide gradients, regardless of changes in depth or  
128 light availability (Fig. 1a,b). Other anoxygenic phototrophs that have been consistently detected  
129 in the chemocline include the PSB *Lamprocystis*, *Thiocystis* and *Thiodictyon* and several GSB of  
130 the genus *Chlorobium* (Tonolla *et al.*, 1999, 2004, 2005). Together these bacteria constituted the  
131 majority of the total phototrophic cells ( $10^6 \cdot \text{ml}^{-1}$ ) in 2015, but they are considerably smaller than  
132 *Chr. okenii*.

133 The oxidation of sulfide by these anoxygenic phototrophs proceeds via the formation of  $\text{S}^0$  as an  
134 obligate intermediate (Mas and Van Gemerden, 1995). This  $\text{S}^0$  was measured as particulate sulfur  
135 on  $0.7 \mu\text{m}$  filters and may comprise  $\text{S}^0$  stored intracellularly by PSB and  $\text{S}^0$  adhering extracellularly  
136 to GSB. The highest concentrations of  $\text{S}^0$  (up to  $45 \mu\text{M}$ ; Fig. 1c) coincided with the highest *Chr.*  
137 *okenii* cell numbers (Fig. 1b) in the chemocline. It is likely that this  $\text{S}^0$  was present in the form of

138 both elemental S and polysulfides formed by the reaction of free sulfide with intra- and  
139 extracellular S<sup>0</sup>, as has previously been suggested in other euxinic lakes (Overmann, 1997). Our  
140 analytical method for total S<sup>0</sup> did not distinguish between different forms of S<sup>0</sup> such as  
141 cyclooctasulfur and polysulfides. However, we could confirm the presence of polysulfides inside  
142 live *Chr. okenii* cells in environmental samples using Raman spectroscopy. The Raman spectrum  
143 of a sulfur inclusion from *Chr. okenii* exhibited two weak peaks at 152 and 218 and a prominent  
144 peak at 462 cm<sup>-1</sup> (Fig. S3) which is characteristic of linear polysulfide species (Janz *et al.*, 1976).  
145 The Raman peak at ~2900 cm<sup>-1</sup> corresponds to the CH<sub>2</sub> and CH<sub>3</sub> stretching vibrations (Socrates,  
146 2004), and its co-occurrence with polysulfide peaks support the theory that the sulfur chains in  
147 these purple sulfur bacteria are terminated by organic end groups as reported previously (Prange  
148 *et al.*, 1999).

149 Over two diurnal cycles, the S<sup>0</sup> inventory (Fig. S4a), or the total amount of particulate S<sup>0</sup> in the  
150 chemocline, was much lower than expected from the sulfide gradients and corresponding sulfide  
151 fluxes (discussed below), suggesting that stored S<sup>0</sup> served only as a transient intermediate and  
152 was rapidly oxidized to sulfate. No day-night trends in S<sup>0</sup> accumulation were apparent in the  
153 chemocline. Nevertheless, the increase in the S<sup>0</sup> inventory, at several time points during the night  
154 was indicative of dark sulfide oxidation.

155 In culture, *Chromatium* spp. are known to store carbon compounds like glycogen and  
156 polyhydroxyalkanoates (PHAs) which have been proposed to be involved in dark sulfur  
157 metabolism (Mas and van Gemerden, 1995). We therefore quantified glycogen and PHA  
158 abundance in biomass samples from one day/night profile of the chemocline (Fig. 2). We could  
159 not detect any PHA, but the presence of glycogen during the day and night coincided with *Chr.*

160 *okenii* cell numbers (Fig. 2). This is consistent with previous reports of glycogen storage and an  
161 absence of PHA in natural populations of *Chr. okenii* (Del Don *et al.*, 1994). While the highest  
162 potential cellular glycogen content ( $2.38 \cdot 10^{-6}$   $\mu\text{g}/\text{cell}$ ) was found at the top of the chemocline  
163 during the day, we observed little change in the cellular glycogen content between day and night  
164 (Fig. S5). Average potential cellular glycogen decreased from  $5.50 \cdot 10^{-7}$   $\mu\text{g}/\text{cell}$  during the day to  
165  $5.33 \cdot 10^{-7}$   $\mu\text{g}/\text{cell}$  during the night, which represents a 3% reduction in cellular glycogen reserves.  
166 This is in contrast with a previous study of storage compounds in natural populations of *Chr. okenii*  
167 in Lake Cadagno which reported 50% decrease in glycogen reserves in the dark (Del Don *et al.*,  
168 1994). This apparent decrease in glycogen reported previously may be a result of undersampling,  
169 as our time- and depth-resolved biogeochemical profiles revealed light-dark independent  
170 variations in *Chr. okenii* cell numbers and glycogen concentrations. While it has been  
171 demonstrated that *Chromatium* sp. in pure cultures obtain energy from the reduction of  $\text{S}^0$  with  
172 glycogen in the dark (Van Gemerden, 1968), we could not confirm this observation for *Chr. okenii*  
173 *in situ*. From our data, we conclude that storage compounds did not play a significant role in the  
174 dark respiratory metabolism of *Chr. okenii* in the Lake Cadagno chemocline.

175 Sulfate was measured as the end product of sulfide oxidation, but due to the high (1-2 mM)  
176 background sulfate concentrations, the comparably small concentration changes resulting from  
177 sulfide oxidation processes are non-detectable. To identify regions of sulfate production in and  
178 around the chemocline, we therefore determined deviations from the sulfate-conductivity mixing  
179 line drawn for each profile (see Fig. S6 for details). Strong mixing of the chemocline is expected  
180 to produce a linear relationship between sulfate and conductivity, and large digressions from this  
181 best-fit line indicated that sulfate was produced faster than the rate of mixing. The expected



182 sulfate concentration could be extrapolated based on measured conductivity, and then  
183 subtracted from the measured sulfate concentration to give excess sulfate:

$$184 \quad \text{measured } [\text{SO}_4^{2-}] - \text{expected } [\text{SO}_4^{2-}] = \text{excess } [\text{SO}_4^{2-}]$$

185 This excess sulfate was attributed to biological sulfate production. Sulfate profiles from 2015  
186 plotted over two diurnal cycles exhibited a peak at the top of the chemocline in the region of  
187 oxygen depletion (Fig. 1d). Interestingly, sulfate production was observed both during and at the  
188 end of the night. The overlap of excess sulfate and oxygen in 2015 profiles was the first indication  
189 that sulfide may be oxidized aerobically, without sunlight. Daytime sulfate production in 2014  
190 related to photosynthetically active radiation (PAR) intensity (Fig. S2), suggesting that sulfide and  
191  $\text{S}^0$  could either have been oxidized aerobically within the chemocline using *in situ*-produced  
192 oxygen (Milucka *et al.*, 2015) or phototrophically. The comparatively broad biogenic sulfate peak  
193 in the 2014 night profile likely reflects the broader vertical distribution of the *Chr. okenii*  
194 population (Fig. S2).

195 The sulfate excess in the chemocline is not expected to be affected by sulfate reduction as no  
196 sulfate reduction was detected within the chemocline in 2014 or 2015. The sulfate reduction rates  
197 measured in the sulfidic zone 1 m below the chemocline were about  $235 \text{ nM}\cdot\text{d}^{-1}$  and  $375 \text{ nM}\cdot\text{d}^{-1}$   
198 in 2014 and 2015, respectively.

199 To quantify biological sulfide consumption over time, we calculated the total sulfide flux into the  
200 chemocline (Fig. S4b). Assuming that phototrophic sulfide oxidation ceases in the dark, upwards-  
201 diffusing sulfide should accumulate in the chemocline at night. The expected sulfide accumulation  
202 was calculated based on fluxes into the layer over a 10-h night period and compared to the actual

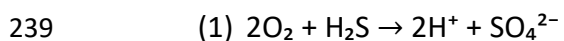
203 sulfide concentration observed in the layer. From an average sulfide flux  $F = 0.15 \mu\text{m}\cdot\text{cm}^{-2}\text{h}^{-1}$  (Fig.  
204 S4b), into a well-mixed layer of thickness  $H = 1 \text{ m}$  over  $t = 10 \text{ hours}$ , the resulting sulfide  
205 concentration  $C = F*t/H$  should be about  $15 \mu\text{M}$  in the chemocline. However, the sulfide  
206 measured in the layer was about  $3 \mu\text{M}$  (Fig. 1a), or five times less, indicating that sulfide is  
207 consumed.

208 We therefore partitioned the total sulfide flux into two fractions: the flux of biologically consumed  
209 sulfide and the flux of residual sulfide in the chemocline. First, the amount of residual sulfide was  
210 calculated at each sampling time point by integrating sulfide concentrations within the mixed  
211 layer (Fig. S4c). The rate of sulfide accumulation was then calculated for each 4-h sampling  
212 interval and subtracted from the total sulfide flux to give the biologically consumed sulfide flux.  
213 The flux of sulfide consumed in the dark was in the same range as in the day ( $0.03$  to  $0.22$   
214  $\mu\text{mol}\cdot\text{cm}^{-2}\text{h}^{-1}$ ) and the residual sulfide flux was very small in comparison (Fig. 3a). The observed  
215 variations did not correlate with day-night cycles and the changes of sulfide gradients could have  
216 been induced by internal waves, as mentioned above. Together, this indicates that sulfide  
217 oxidation continued in the dark and seemed to be related to the total sulfide flux (Fig. S4b) rather  
218 than the presence of sunlight. For comparison, the upwards flux of sulfide in previous years was  
219 slightly lower, or  $0.011$ - $0.024 \mu\text{mol}\cdot\text{cm}^{-2}\text{h}^{-1}$  in 2013 and  $0.032$ - $0.072 \mu\text{mol}\cdot\text{cm}^{-2}\text{h}^{-1}$  in 2014.

220 It was not possible to calculate  $S^0$  fluxes in Lake Cadagno because  $S^0$  is actively transported by the  
221 motile purple sulfur bacteria during chemo- and phototaxis (Pfennig *et al.*, 1968) independent of  
222 diffusive processes. The total (upwards and downwards) biogenic sulfate flux (Fig. S4d) in this  
223 region was roughly equivalent to the sulfide flux and followed a similar trend.

224 Overall, our high-resolution profiles revealed that sulfide in Lake Cadagno was consumed during  
225 the day and night, but only light-dependent sulfide oxidation has thus far been recognized as a  
226 major sulfide-removing process in the lake. In the absence of light, it is also possible that  
227 alternative electron acceptors such as  $\text{NO}_x^-$ ,  $\text{Fe}^{3+}$ ,  $\text{Mn}^{4+}$ , or  $\text{O}_2$  play a role in sulfide oxidation.  
228 Nitrate and nitrite concentrations in the Lake Cadagno chemocline are negligible (Halm *et al.*,  
229 2009; Milucka *et al.*, 2015). High fluxes of reduced, dissolved metals ( $0.027 \mu\text{mol Fe}\cdot\text{cm}^{-2}\cdot\text{d}^{-1}$  and  
230  $\text{Mn } 0.049 \mu\text{mol Mn}\cdot\text{cm}^{-2}\cdot\text{d}^{-1}$ ) suggest that Fe- and Mn-oxides are rapidly reduced by  
231 microorganisms or abiotically by sulfide in the chemocline (Berg *et al.*, 2016), but re-oxidation of  
232 Fe and Mn would ultimately depend on oxygen in the dark. We therefore considered oxygen as  
233 the principal direct (or indirect) oxidant responsible for observed dark sulfide oxidation.

234 The oxygen flux into the chemocline varied slightly between  $0.022\text{-}0.071 \mu\text{mol}\cdot\text{cm}^{-2}\cdot\text{h}^{-1}$  over the  
235 period of 48 h (Fig. 3b). Oxygen fluxes measured in 2013 and 2014 were in the same range, or  
236  $0.013\text{-}0.048 \mu\text{mol}\cdot\text{cm}^{-2}\cdot\text{h}^{-1}$  and  $0.037\text{-}0.073 \mu\text{mol}\cdot\text{cm}^{-2}\cdot\text{h}^{-1}$ , respectively. To relate oxygen fluxes to  
237 sulfide consumption, we assumed a 2:1 stoichiometry between oxygen and sulfide for aerobic  
238 sulfide oxidation to sulfate:



240 If all oxygen was used to respire sulfide, calculated oxygen fluxes in 2013 and 2014 were in all  
241 cases sufficient to account for the sulfide oxidized in the dark. In 2015, aerobic sulfide respiration  
242 could account for up to 10-50% of sulfide oxidized during the day and 5-45% of sulfide oxidized  
243 during the night (Fig. 3c). During the day, the remainder of sulfide oxidation could be attributed  
244 to anoxygenic photosynthesis and/or aerobic sulfide oxidation fueled by *in situ* oxygen production

245 by photosynthetic algae. At several time points in the dark, however, we could not explain the  
246 disappearance of roughly 60-90% of upwards-diffusing sulfide. We hypothesize that the missing  
247 oxygen is supplied laterally from the turbulent transport initiated by internal wave breaking at  
248 the lake boundaries. The convection within the chemocline may be key to the transport of oxygen  
249 and sulfide to aerobic sulfide-oxidizing bacteria in the chemocline. A weakening of the mixing  
250 regime was observed in August 2015 (Sommer *et al.*, 2017) which may have signified a slowed  
251 transport of electron acceptors, thus contributing to the accumulation of sulfide in the  
252 chemocline.

### 253 ***Mixing and bacterial motility in Lake Cadagno***

254 To test the importance of lateral and vertical mixing, we set up simplified laboratory incubations  
255 where water from Lake Cadagno chemocline was inoculated into agar-stabilized sulfide gradient  
256 tubes. After five weeks of incubation under permanent light conditions, dense communities of  
257 PSB developed between the gradient of upwards-diffusing sulfide and the surface colonies of  
258 photosynthetic algae (Fig. S7). Microsensor profiles revealed that sulfide was completely  
259 consumed at the base of the PSB layer in the light, but as soon as the light was turned off, the  
260 sulfide gradient diffused upwards through the agar into the zone of purple bacteria. This is in  
261 contrast to the sulfide profiles in the lake where irrespective of the day-night cycle, sulfide is  
262 consistently consumed at the bottom of the chemocline. We speculate that restricted bacterial  
263 motility in the agar and diffusion-limited conditions may have accounted for the differences  
264 observed between our cultures and *in situ* sulfide consumption as bacterial motility and mixing  
265 conditions appear necessary for continued dark sulfide oxidation in Lake Cadagno.

266 In fact, we could confirm that *Chr. okenii* are highly motile both in the day and the night by  
267 performing dark field video microscopy (see Movie S1 in Supplementary Materials) of  
268 environmental samples obtained during the night and monitored in a dark room to avoid any  
269 light-induced artefacts. Although the average night time swimming speed of *Chr. okenii* ( $9.9 \mu\text{m}$   
270  $\text{s}^{-1}$ ; see Fig. S8) was a third of the day time swimming speed ( $27 \mu\text{m} \text{s}^{-1}$ ; Sommer et al. 2017), it is  
271 clear that *Chr. okenii* remains motile even under dark conditions.

### 272 ***Metagenomic insights into the Chromatium okenii population in Lake Cadagno***

273 To assess whether the genomic potential supports light-independent, aerobic sulfide oxidation  
274 by *Chr. okenii* in Lake Cadagno, we sequenced two metagenomes, one from the Lake Cadagno  
275 chemocline and one from the phototrophic, sulfide-oxidizing enrichment culture in an agar tube  
276 described above (Table S1). From a combined metagenomics assembly, we reconstructed a high  
277 quality (90% complete, <1% contaminated) metagenome-assembled genome (MAG) of a PSB  
278 highly abundant in the sulfur-oxidizing enrichment culture (Fig. S9). The recovered MAG had a  
279 low average nucleotide identity ANI (<70%) to any sequenced *Chromatiaceae* genomes (data not  
280 shown). However, it encoded an rRNA operon, including a complete 16S rRNA gene with 99%  
281 sequence identity to the 16S rRNA gene of *Chr. okenii* (Imhoff et al., 1998; Tonolla et al., 1999),  
282 and thus likely represents a strain of *Chr. okenii* which is the type strain of the genus *Chromatium*.  
283 At this time, *Chr. okenii* has not been successfully isolated in pure culture, nor is there any  
284 published genome available for this organism.

285 The key metabolic process of *Chr. okenii* in Lake Cadagno is photoautotrophic sulfur oxidation. In  
286 accordance, the *Chr. okenii* MAG contained genes encoding for a sulfide : quinone reductase (*sqr*)

287 and the full genomic inventory encoding for a reverse-acting dissimilatory sulfite reductase (rDSR)  
288 pathway (Fig. 4). The operon structure of the rDSR encoding genes (*dsrABEFHCMKLJOPN*) was  
289 identical to the operon structure in the well described PSB model organism *Allochromatium*  
290 *vinosum* (Dahl *et al.*, 2005), but no *dsrR* and *dsrS* gene were found. No genes encoding for sulfur  
291 oxidation via the SOX pathway, or homologues of sulfur globule proteins (*sgpABC*) typically found  
292 in PSB were detected in the draft genome. In line with its phototrophic metabolism, the *Chr.*  
293 *okenii* MAG showed the genomic potential for photosynthesis, with the genes encoding for a light  
294 harvesting complex 1 (*pufAB*) and a PSB-type photosynthetic reaction center (*pufLMC*) encoded  
295 in a single operon. Furthermore, the full genomic repertoire for a NADP-Me type C4  
296 photosynthetic carbon assimilation cycle, and all genes (with exception of *cbbS* encoding for the  
297 small subunit of the ribulose-1,5-bisphosphate carboxylase/oxygenase) necessary for CO<sub>2</sub>  
298 assimilation via the Calvin-Benson-Bassham (CBB) Cycle were present (Fig. 4).

299 Many *Chromatiaceae* can grow chemoautotrophically, respiring oxygen under microoxic  
300 conditions (Kämpf and Pfennig, 1980). Cytochrome (Cyt) *c*-containing oxidases (e.g. Cyt *aa3*, Cyt  
301 *cbb3*) were not found in the *Chr. okenii* MAG. However, a Cyt *bd* type ubiquinol oxidase, known to  
302 function as sulfide-resistant O<sub>2</sub>-accepting oxidase in other *Gammaproteobacteria* (Forte *et al.*,  
303 2016), was identified (Fig. 4). Further, a plethora of genes related to heme *b* (*gltX*, *hemALBCD*,  
304 and *hemH*) and siroheme (*cysG*) synthesis, degradation (a heme oxygenase) and export (ABC-type  
305 heme exporter, *ccmABCD*), as well as hemerythrin-like metal binding proteins were encoded.  
306 Hemerythrin has been implicated in binding of oxygen for delivery to oxygen-requiring enzymes,  
307 for detoxification, or for oxygen sensing in motile, microaerobic prokaryotes (French *et al.*, 2007).  
308 The presence of these oxygen-dependent enzymes, as well as a key oxidative stress defense

309 enzyme superoxide dismutase (SOD), support the idea that *Chr. okenii* may be facultatively  
310 microaerobic. A complete set of genes for flagellar biosynthesis (*fliDEGHJKLMNQPQRW*,  
311 *flgABCDEFGHIK*, *flhAB*) and flagellar motor proteins (*motAB*) confer motility to this bacterium.  
312 Several other genes revealed interesting metabolic capacities of *Chr. okenii*. A cytosolic  
313 bidirectional [NiFe] type 3d hydrogenase and a nitrogenase were encoded in the MAG (Fig. 4),  
314 implicating the potential for involvement of *Chr. okenii* in nitrogen fixation and hydrogen  
315 oxidation which has previously been overlooked. Additionally, the *Chr. okenii* MAG encoded a  
316 glycogen synthase and a glycogen debranching enzyme, as well as the full genomic repertoire  
317 necessary for polyhydroxyalkanoate (PHA) biosynthesis. This is consistent with the detection of  
318 glycogen in our biogeochemical profiles of the chemocline. Finally, it is possible that novel  
319 terminal oxidases are among the hypothetical genes that could not be assigned any known  
320 function.

## 321 **CONCLUSIONS**

322 It is intriguing that oxygen should play a major role in sulfide oxidation in the ostensibly anoxic  
323 chemocline of Lake Cadagno, especially by purple sulfur bacteria generally thought to lead an  
324 anaerobic lifestyle. To explain the coupling of oxygen and sulfide consumption in the oxygen- and  
325 sulfide-free chemocline of Lake Cadagno, we sketched a diagram of the transport processes likely  
326 driving biological activity in the chemocline (Fig. 5). As described in Sommer *et al.*, (2017), active  
327 convection of the chemocline can be driven by the formation of sinking bacterial plumes.  
328 Combined with turbulence induced by the breaking of internal waves at sides of the lake basin,  
329 these convective currents may entrain sulfide and oxygen at the boundaries of the chemocline  
330 and fuel populations of sulfide-oxidizing *Chr. okenii* there.

331 Sulfur-oxidizing bacteria have previously been reported to bridge distances between pools of  
332 electron donors and acceptors by intracellularly storing and transporting  $S^0$  and  $NO_3^-$  between  
333 redox zones (Fossing *et al.*, 1995; Jørgensen and Gallardo, 1999) and even by transferring  
334 electrons along nanowires (Pfeffer *et al.*, 2012), but the sulfide oxidation processes in Lake  
335 Cadagno represent a new mechanism of electron acceptor/donor coupling across large distances.  
336 After entrainment into the chemocline, dissolved oxygen and sulfide are consumed so rapidly that  
337 they remain below detection limits. Our metagenomic evidence shows that *Chr. okenii* possesses  
338 several high-affinity oxidases which may enable it to respire oxygen at such low, nanomolar  
339 concentrations. The physical and biological processes described here may therefore provide clues  
340 to sulfide oxidation in other anoxic environments such as the Black Sea where the mechanism of  
341 sulfide removal is not completely understood. Clearly, the biochemical limits to oxygen utilization  
342 are far below current definitions of *anoxia* and demonstrate that aerobic respiration is possible  
343 in so-called “anoxic” lacustrine (Milucka *et al.*, 2015) and marine (Garcia-Robledo *et al.*, 2017)  
344 waters.

345 Overall, we show that in contrast to observations from laboratory cultures, *Chr. okenii* appear to  
346 have a very different metabolism in the environment where high fluxes of nutrients rather than  
347 absolute nutrient concentrations fuel microbial activity. The unexpected insights into the  
348 ecophysiology of the purple sulfur bacteria obtained here demonstrate the importance of  
349 studying these versatile bacteria *in situ* using culture-independent methods to understand their  
350 environmental function.

## 351 **EXPERIMENTAL PROCEDURES**

### 352 **Sampling**



353 The meromictic Lake Cadagno is situated in the Piora Valley in the Swiss Alps at an altitude of  
354 1921 m. Data presented here were collected during field campaigns in September 2013, August  
355 2014, June 2015 and August 2015. In 2013 and 2014 *in situ* measurements were performed with  
356 a profiling ion analyzer (PIA; see Kirf *et al.*, 2014 for description) lowered from a platform  
357 anchored at the deepest part of the lake (20.7 m). Conductivity, turbidity, depth (pressure),  
358 temperature and pH were measured with a multi-parameter probe (XRX 620, RBR). Dissolved  
359 oxygen was recorded online with a type PSt1 normal (detection limit 125 nM) micro-optode and  
360 a type TOS7 trace (reliable detection limit 50-100 nM) micro-optode (PreSens). The oxygen  
361 sensors were calibrated by parallel Winkler titrations. Water samples for chemical analyses and  
362 cell counts were collected with a rosette syringe sampler equipped with twelve 60-ml syringes  
363 triggered online at selected depths. Due to a technical failure of the PIA, the 6 AM profile in August  
364 2014 and all subsequent profiles in 2015 were measured with a SBE 19 plus V2 CTD probe (Sea-  
365 Bird Electronics, WA, USA) equipped with sensors for pressure, temperature and conductivity,  
366 and with additional sensors for turbidity (WET Labs Eco), oxygen (SBE 43), pH (18-I) and two  
367 fluorescence wavelengths (WET Labs ECO-AFL, FL, USA). The detection limit of the SBE 43 oxygen  
368 probe was about 1  $\mu\text{mol/l}$ . In parallel with *in situ* measurements, water for chemical analyses was  
369 pumped to the surface through Neoprene tubing attached to the CTD and filled into 60-ml  
370 syringes (rinsed 2 X with *in situ* water) on board. Two parallel metal plates of diameter  $\sim 15$  cm  
371 attached to the submersed end of the tubing served to channel water horizontally, resulting in  
372 more discrete vertical profiling.

373 Water samples in syringes were aliquoted on board immediately after collection. Samples for  
374 sulfate analyses were filtered (0.22  $\mu\text{m}$  pore size) directly into sterile Eppendorf vials. Sulfide

375 samples were fixed with Zn-acetate to a final concentration of 0.1 % (w/v). Biomass was  
376 concentrated onto glass fiber filters (0.7  $\mu\text{m}$  pore size) and stored at  $-20^{\circ}\text{C}$  for analyses of  
377 intracellularly stored elemental sulfur and organic carbon compounds. Filtrate (0.22  $\mu\text{m}$  pore size)  
378 was also collected and frozen at  $-20^{\circ}\text{C}$  for metabolome analysis of dissolved compounds. Samples  
379 for fluorescence *in situ* hybridization were immediately fixed with 2% (v/v) formaldehyde.  
380 Samples for DNA analysis were collected from the chemocline in August 2014 by concentrating  
381 microbial cells on polycarbonate filters (0.22  $\mu\text{m}$  pore size) on site and freezing at  $-20^{\circ}\text{C}$  until  
382 further processing.

383 Additional water for cultivation and motility experiments was pumped directly from the  
384 chemocline into 1-L Duran bottles and sealed with butyl rubber stoppers without a headspace to  
385 maintain anoxic conditions.

### 386 **Chemical Analyses**

387 Sulfide was measured using the colorimetric method of Cline (1969). Sulfate was measured on a  
388 761 Compact ion chromatograph (Metrohm, Filderstadt, Germany) equipped with a Metrosep A  
389 SUPP 5 column. Intracellular sulfur on filters was extracted by sonication in methanol for 15 min  
390 in an ice bath. Samples were analyzed on an Acquity H-Class UPLC system (Waters Corporation,  
391 USA) with an Acquity UPLC BEH C18 column coupled to a photodiode array (PDA) detector using  
392 UPLC-grade methanol as eluent. Data was acquired and processed using the Empower III  
393 software.

394 Intracellular glycogen was analyzed following the procedures of the assay kit (MAK016 Sigma  
395 Aldrich). Briefly, cells were extracted by scraping them from GFF filters and homogenizing in 200

396  $\mu\text{L}$  extraction buffer and centrifuged two times to clear the supernatant. The supernatant was  
397 analyzed fluorometrically after incubation with enzyme mix and fluorescent peroxidase substrate.  
398 Intracellular PHA was analyzed using the protocol from Braunegg *et al.* (1978). Hydrolyzation of  
399 the polymer and conversion to a methyl-ester of the monomeric hydroxyalkanoate fraction was  
400 done in acidified alcohol solution (6%  $\text{H}_2\text{SO}_4$  in methanol) and chloroform under heating ( $100^\circ\text{C}$ ,  
401 2h). After addition of water and phase separation the organic phase was analyzed with GC-MS  
402 (Agilent 7890B GC connected to Agilent 5977A MSD) to detect the methylhydroxyalkanoates  
403 using the following settings: Agilent 30 m DB-5-MS column, splitless injection of  $1\ \mu\text{L}$ , temperature  
404 program was  $50^\circ\text{C}$  for 1min than heating  $10^\circ\text{C}/\text{min}$  until  $120^\circ\text{C}$  followed by  $45^\circ\text{C}/\text{min}$  until  $320^\circ\text{C}$   
405 and hold for 5 minutes. Benzoic acid was used as internal standard in each sample and  
406 quantification was done with pure polyhydroxybutyrate standard (Sigma\_Aldrich).

407 Sulfate reduction rates were measured by adding the radiotracer  $^{35}\text{SO}_4^{2-}$  (5 MBq) to anoxic lake  
408 water in 50-ml glass syringes and incubated in the dark. A solution of unlabeled  $\text{Na}_2\text{S}$  was added  
409 to a final concentration of  $\sim 50\ \mu\text{mol}\cdot\text{l}^{-1}$  as a background sulfide pool in case of sulfide re-oxidation.  
410 At each sampling point, 10 ml of sample was dispensed into 5 ml of 20% (w/v) Zn-acetate.  
411 Reduced sulfur species (e.g. sulfur and sulfide as ZnS) were separated out via the chromium  
412 distillation method described in (Kallmeyer *et al.*, 2004) and the radioactivity per sample was  
413 determined via scintillation counting (Packard 2500 TR).

#### 414 **Confocal Raman spectroscopy**

415 In glove box under 90:10  $\text{N}_2\text{-CO}_2$  atmosphere, a drop of fresh sample from the chemocline was  
416 mounted between two glass coverslips and sealed with electrical tape to prevent contact with

417 air. A polysulfide solution containing 5.06 g  $\text{Na}_2\text{S} \cdot 9\text{H}_2\text{O}$  and 5.8 g elemental sulfur per 100 ml  
418  $\text{H}_2\text{O}$ , with a final pH of 9.5 and sulfide concentration of 210 mM was used as reference.  
419 Measurements were conducted with an NTEGRA Spectra confocal spectrometer (NT-MDT,  
420 Eindhoven, Netherlands) coupled to an inverted Olympus IX71 microscope. The excitation light  
421 from a 532-nm solid-state laser was focused on the sample through an Olympus 100X (numerical  
422 aperture [NA], 1.3) oil immersion objective. Raman scattered light was collected by an electron-  
423 multiplying charge-coupled device (EMCCD) camera (Andor Technology, Belfast, Northern  
424 Ireland) cooled to  $-70^\circ\text{C}$ . Spectra were recorded between 0 and  $4,500\text{ cm}^{-1}$  with a spectral  
425 resolution of  $0.2\text{ cm}^{-1}$  and analyzed with the software NT-MDT software Nova\_Px 3.1.0.0.

#### 426 **Flux Calculations**

427 Turbulent fluxes ( $J$ ) of sulfide, sulfur, sulfate, and oxygen at the chemocline were calculated  
428 assuming steady state by applying Fick's first law:  $J = -D\partial C/\partial x$ . For sulfide, sulfate, and oxygen we  
429 used the turbulent diffusion coefficient ( $D$ ) of  $1.6 \times 10^{-6}\text{ m}^2\text{ s}^{-1}$  from (Wüest, 1994) corresponding  
430 to turbulence at the Lake Cadagno chemocline boundaries. For sulfur gradients within the well-  
431 mixed chemocline the coefficient  $D = 1.5 \times 10^{-5}\text{ m}^2\text{ s}^{-1}$  (Wüest, 1994) was used. The change in  
432 concentration ( $\partial C$ ) was computed for each species over the depths with the steepest gradients.  
433 Oxygen and sulfide fluxes were determined for the regions immediately above and below the  
434 chemocline, defined as the zone of constant conductivity.

#### 435 **Microbial cultivation**

436 Anoxygenic phototrophic bacteria from the Lake Cadagno chemocline were cultivated in agar-  
437 stabilized, sulfide gradient medium in anoxic test tubes. Solid agar (1.5% w/v agar) and semi-solid  
438 agar (0.25% w/v agar) were prepared separately by autoclaving triple-washed agarose and sterile-

439 filtered water from the Lake Cadagno chemocline, and degassing for 1 h with mixture of 80% N<sub>2</sub>  
440 and 20% CO<sub>2</sub> during cooling to ~50°C. The solid agar was amended with a sterile Na<sub>2</sub>S solution to  
441 a final concentration of ~4mM before pouring into degassed test tubes to form a ~2 cm bottom  
442 layer and allowed to set. The semisolid agar was amended with vitamins and trace elements as  
443 described for cultivation of purple sulfur bacteria (Eichler and Pfennig, 1988) before pouring a ~7  
444 cm top layer, and immediately capped with a butyl rubber stopper. After cooling to ~30°C, 1 ml  
445 of fresh Cadagno chemocline water was used to inoculate the top agar via a degassed syringe.  
446 Tubes were inverted once to mix and allowed to set. Agar cultures were incubated under low, 24-  
447 h light conditions at 15°C to favor the development of anoxygenic phototrophs.

#### 448 **Microsensor measurements**

449 Gradients of pH and H<sub>2</sub>S in agar cultures were measured using microelectrodes built in-house as  
450 described previously (Jeroschewski *et al.*, 1996; de Beer *et al.*, 1997). Immediately before use, the  
451 pH sensor was calibrated in standard buffers and the H<sub>2</sub>S sensor was calibrated in a dilution series  
452 of an acidified Na<sub>2</sub>S solution. Electrodes were mounted on a micromanipulator connected to a  
453 computer and profiles were measured in 250 µm intervals from the agar surface to the base of  
454 the sulfide plug. Agar tubes were uncapped for the insertion of microsensors, and the headspace  
455 was flushed with N<sub>2</sub> gas before recapping immediately after each measurement. Total sulfide  
456 concentrations were calculated from pH and H<sub>2</sub>S gradients as described in Schwedt *et al.*, (2012).

#### 457 **Motility analysis**

458 Water samples containing *Chr. okenii* cells were collected under anoxic conditions from the  
459 chemocline during the night, protected from artificial light with aluminum foil, and analyzed

460 immediately on site. Motile cells were transferred via a degassed glass syringe to a sealed  
461 rectangular millimetric chamber (dimensions 20 mm × 10 mm × 2 mm) prepared using glass slides  
462 separated by a 2-mm thick spacer, which provided an anoxic environment during motility  
463 characterization. Experiments were conducted in a dark room, and imaging was performed using  
464 the dark field microscopy mode at 25 fps, with the lowest intensity illumination. No transient  
465 response was observed right at the start of the imaging, and the swimming velocity remained  
466 steady throughout the duration of the measurements. This is in contrast to swimming behavior  
467 at higher light intensities where the swimming cells exhibited a positive phototactic response  
468 (Sommer *et al.*, 2017). We could therefore rule out a light-induced effect on motility at the  
469 minimum illumination level used for our measurements. Videos of swimming cells were acquired  
470 and subsequently analyzed using the ImageJ Particle Tracker routine to obtain the coordinates of  
471 the cells (geometric centers) at each time interval. These were used to calculate the swimming  
472 speeds and extract the trajectories of individual cells.

### 473 **DNA extraction, sequencing, and analysis**

474 Environmental DNA was extracted from polycarbonate filters with the Ultra Clean MoBio  
475 PowerSoil DNA kit (MoBio Laboratories, Carlsbad, USA) according to the manufacturer's protocol  
476 with the following modification: the bead beating step was reduced to 30 sec followed by  
477 incubation on ice for 30 sec, repeated 4x. The DNA was gel-purified using SYBR Green I Nucleic  
478 Acid Gel Stain (Invitrogen) and the QIAquick Gel Extraction Kit (Qiagen) according to the  
479 accompanying protocols. DNA concentration was determined fluorometrically at 260 nm, using  
480 the Qubit 2.0 Fluorometer and the Qubit dsDNA HS Assay KIT (Invitrogen) and sent to the Max  
481 Planck-Genome Centre (Cologne, Germany) for sequencing. The metagenome was sequenced

482 (100 bp paired end reads) by Illumina HiSeq (Illumina Inc., USA) sequencing following a TruSeq  
483 library preparation. Metagenomic reads were adapter- and quality-trimmed (phred score 15,  
484 bbdduk function of the BBMap package, <https://sourceforge.net/projects/bbmap/>) and paired-  
485 end reads were *de novo* assembled with the uneven depth assembler IDBA-UD (Peng *et al.*, 2012).  
486 The metagenome assembly was binned based on tetranucleotide frequencies, differential  
487 coverage, taxonomic classification, and conserved single-copy gene profiles with the Metawatt  
488 binning software (version 3.5.2; Strous *et al.*, 2012). The completeness and contamination of the  
489 binned MAGs was evaluated with CheckM (Parks *et al.* 2014). The bulk metagenome and the MAG  
490 identified as *Chr. okenii* were automatically annotated in IMG (Markowitz *et al.* 2011), and the *Chr.*  
491 *okenii* MAG was manually screened for the presence of genes of interest to this study. Assembled  
492 data is available in IMG, under the IMG genome IDs 3300010965 (bulk assembly) and 2700988602  
493 (*Chr. okenii* MAG).

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Fig 1: (a) Combined oxygen (top) and sulfide (bottom) profiles of the Lake Cadagno water column revealing the persistence of an oxygen- and sulfide- free zone over a period of 48 hours, with contour lines indicating sulfide concentrations. The bold contour lines delimiting the region with  $> 5 \mu\text{M}$  sulfide were used to define the chemocline in parallel profiles of *Chr. okenii* cell counts (b), particulate  $\text{S}^0$  (c), and sulfate (d). Black dots represent sampling points for all parameters except  $\text{O}_2$  which was measured with a microsensor mounted on a CTD probe. Shaded boxes represent dark periods between sunset at  $\sim 20:50$  and sunrise at  $\sim 6:10$ . Time plots were interpolated from original profiles measured in August 2015 and are provided in Fig S1.

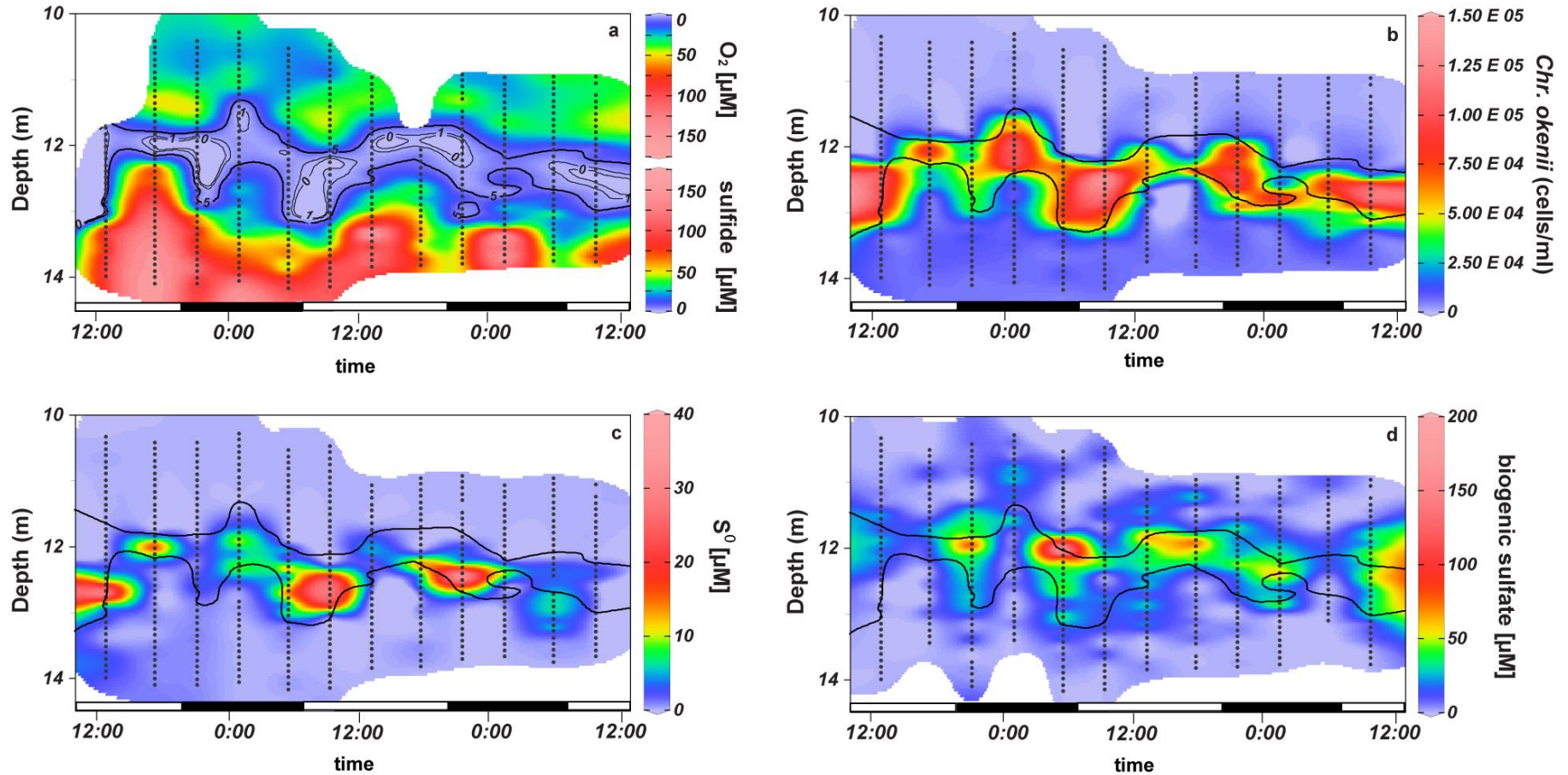


Figure 2: A day (13:00) and a night (1:30) profile through the chemocline illustrating glycogen and  $S^0$  concentrations in relation to *Chr. okenii* cell numbers, oxygen, and sulfide gradients in the chemocline. Profiles were measured in August 2015. PHA was below detection limits and no oxygen data is available for the day profile.

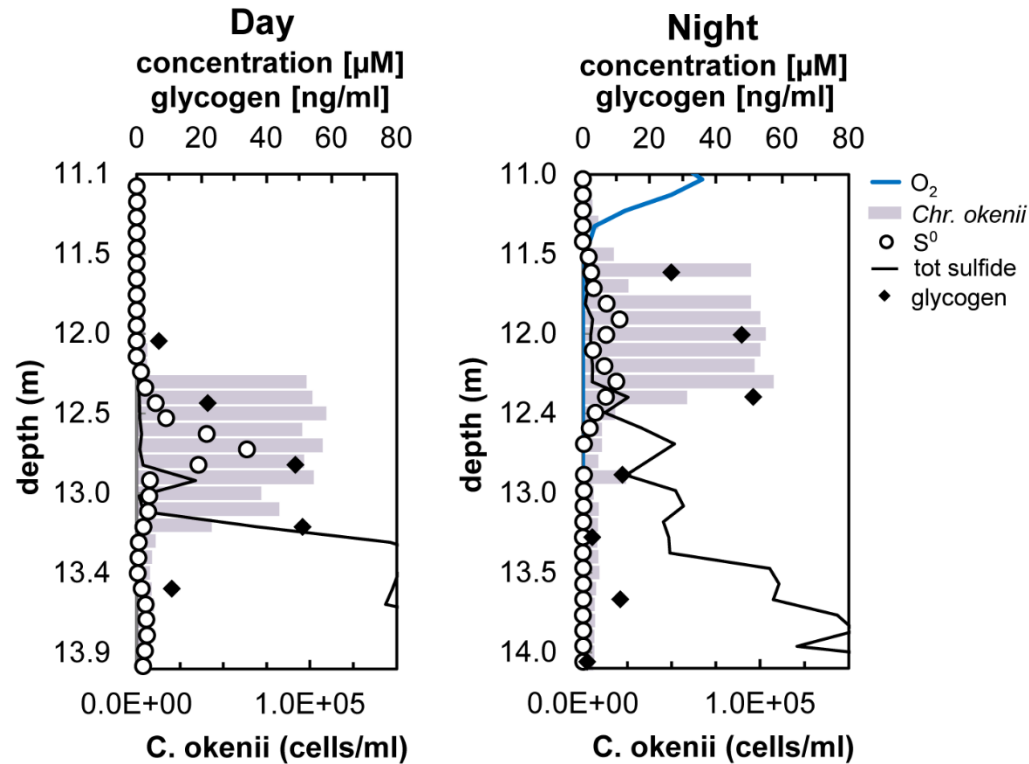


Figure 3: Sulfide and oxygen fluxes in the Lake Cadagno chemocline were calculated from profiles measured 4-h intervals over 2 day-night cycles. (a) The consumed sulfide flux (solid line) was calculated by subtracting the residual sulfide flux (dashed line) from the total sulfide flux into the mixed layer. (b) The downwards oxygen flux into the chemocline was used to estimate (c) the maximum % of sulfide aerobically respired, assuming the complete oxidation of sulfide to sulfate. Shaded regions represent dark periods.

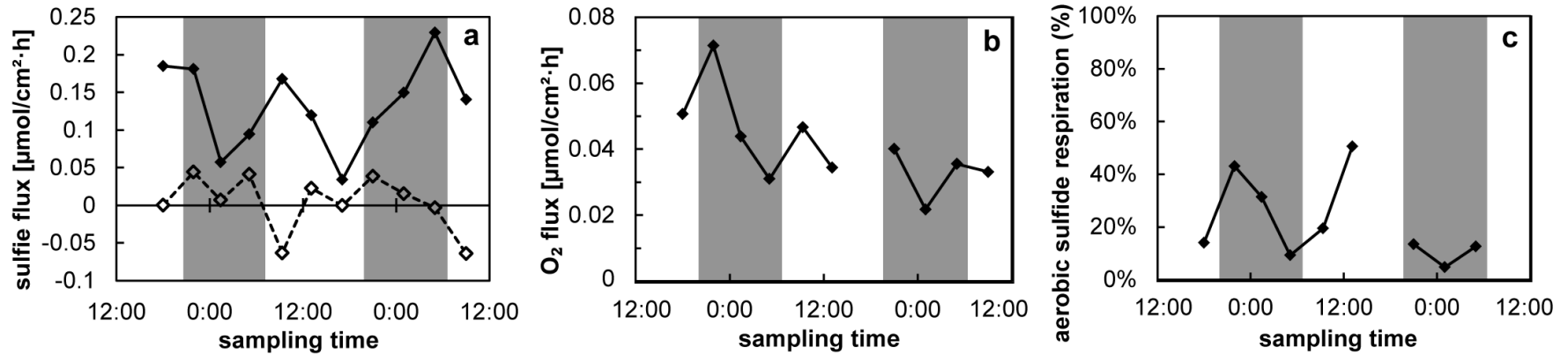


Figure 4: *Chr. okenii* cell illustration, showing the metabolic potential inferred from the metagenome-assembled genome with a particular focus on the genetic machinery implicated in photosynthesis, sulfur oxidation, aerobic metabolism, motility, glycogen and PHA storage, nitrogen fixation and transmembrane transport. The respiratory chain enzyme complexes are labeled with Roman numerals.

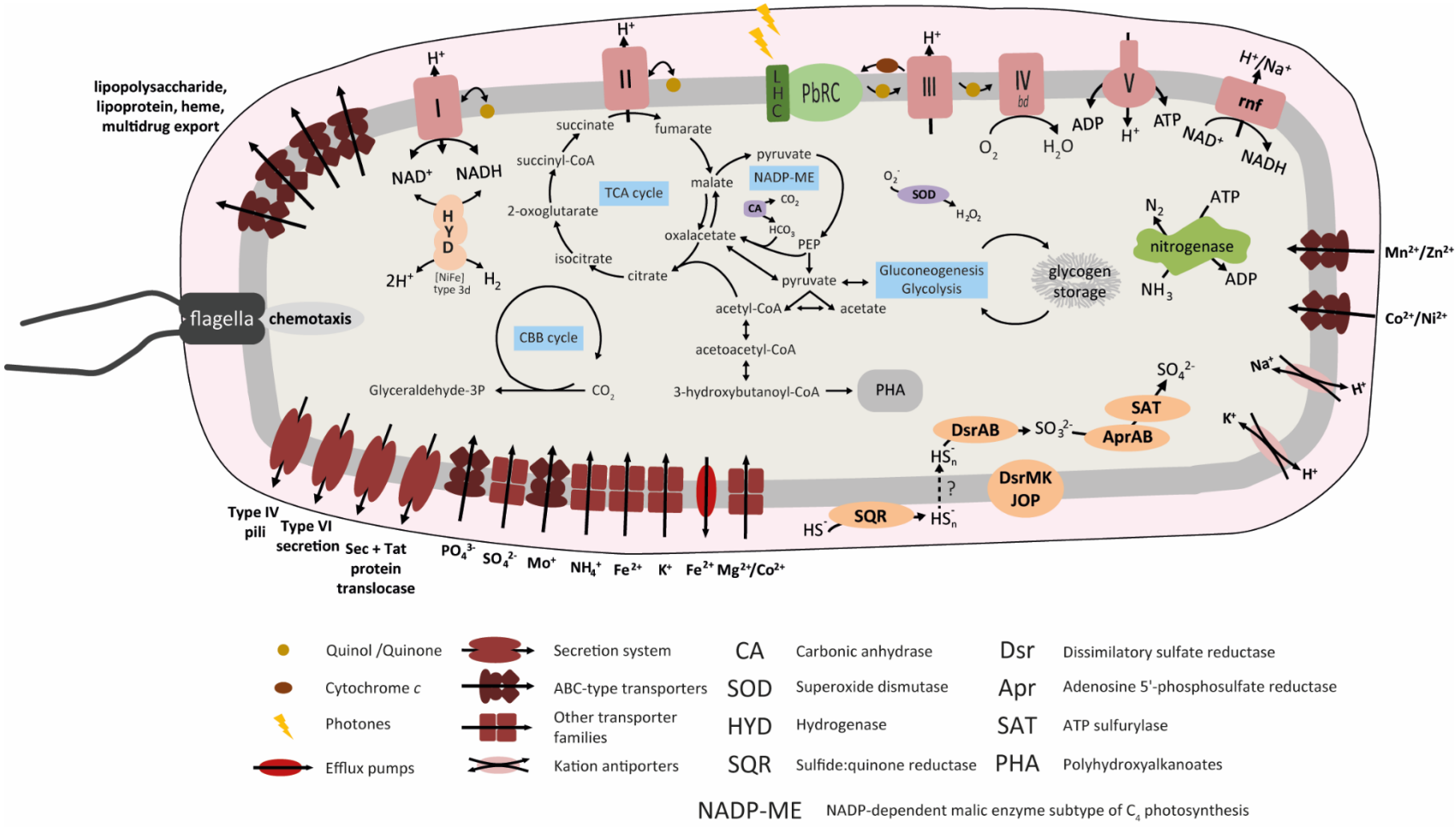


Figure 5: Schematic of phototrophic and aerobic sulfide oxidation processes in the Lake Cadagno chemocline. Convection in the chemocline may be driven by a combination of turbulence and sinking bacterial plumes, represented by the large number of descending *Chr. okenii* cells on the left. As a result, oxygen and sulfide are entrained into the chemocline and immediately consumed by purple sulfur bacteria, keeping concentrations of these compounds below detection limits. *Chr. okenii* cells, depicted with internal sulfur globules (yellow dots), are pulled in the direction of their flagellar bundle.

